

## GEOL 700(i): Geophysical Geodesy - Day 5

### Preliminaries

- Issues with Problem Set #2? Questions? Revelations?
- Flesh et al., 2000, led by Jay.

### New Readings

- STRAIN: Turcotte and Schubert, Segall

Note you can now read Paul Segall's book "Earthquake and Volcano Deformation" online via the UNR library web page.

go to <http://www.knowledgecenter.unr.edu/>

under Start Searching type in name of the book, hit return. Click on the right choice, note it says "Reconstruction German Histories" .... not sure why.... click on "Full-text version (UNR users only)".

### Strain, and Strain Rate

Eulerian vs. Lagrangian descriptions of deformation.

In two-dimensions:

The basics in changes of shape.

$$\epsilon_{xx} = \frac{\partial u_x}{\partial x}, \quad \epsilon_{yy} = \frac{\partial u_y}{\partial y}, \quad \epsilon_{xy} = \frac{\partial u_x}{\partial y} + \frac{\partial u_y}{\partial x}.$$

Tensor shear strain: 
$$\epsilon_{ij} = \frac{1}{2} \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

Engineering shear strain: 
$$\epsilon_{ij} = \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

Two different conventions. Remember to know which one you are using.

Note  $\epsilon_{yx} = \epsilon_{xy}$ .

The strain (rate) tensor  $\boldsymbol{\epsilon} = \begin{bmatrix} \epsilon_{xx} & \epsilon_{yx} \\ \epsilon_{xy} & \epsilon_{yy} \end{bmatrix}$  keeps track of all three.

Here,  $u_x$  is displacement so  $\boldsymbol{\epsilon}$  is strain. Alternatively  $u_x$  could be velocity, in which case  $\boldsymbol{\epsilon}$  would be strain *rate*.

Velocity gradients are not (necessarily) strain rates. For example, on a rotating plate, velocity gradients are not zero, but strain rate is zero. We must separate the displacement tensor into components of strain and rotation.

Given a velocity field that is smoothly varying, in a local patch, we can estimate a constant strain rate tensor value. This can be separated into rotation and strain parts if the network is “strong”.

Examples of strong vs. weak networks.

### Principal strain axes

These are the eigenvectors of the strain tensor i.e. the vectors that satisfy:

$$\boldsymbol{\varepsilon} \mathbf{dx} = A \mathbf{dx}$$

where  $\boldsymbol{\varepsilon}$  is the strain tensor, and  $A$  is a scalar. These two eigenvectors have directions which correspond to the principal strain rate axes. The magnitude of these two vectors are the principal strain rates  $\varepsilon_1$  and  $\varepsilon_2$ .  $\varepsilon_1$  is usually the larger value so  $\varepsilon_1 - \varepsilon_2$  is positive.

### Strain invariants

There are properties of the strain tensor that do not change with changes in coordinate axes (e.g. rotations, and translations).

First Invariant:

$$\varepsilon_{xx} + \varepsilon_{yy} = \text{tr}(\boldsymbol{\varepsilon}) = \varepsilon_1 + \varepsilon_2$$

The first invariant is always the trace of the tensor (in 2D or 3D). It corresponds to the dilatational (volumetric) component of the shape change.

Second Invariant:

formally

$$\varepsilon_{xx} \varepsilon_{yy} - \varepsilon_{xy}^2$$

Another form is commonly used (C. Kreemer personal communication) which has the advantage of being in units of strain (not strain squared):

$$\sqrt{\varepsilon_{xx}^2 + \varepsilon_{yy}^2 + 2\varepsilon_{xy}^2} \quad (\text{sometimes shown squared})$$

This contains shear and dilatation components of deformation.

Also the difference note that in terms of principal strains:

$$\varepsilon_1 - \varepsilon_2$$

is invariant to coordinate axis changes. It represents shear strain because if the principal strains are equal (isotropic dilatation) then the shear is zero. If  $\varepsilon_1 = -\varepsilon_2$  then dilatation is zero and shear is  $2\varepsilon_1$ .